



On the equivalent Biot and Skempton coefficients of fractured rocks and their impact on the HM behavior of geological media

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Biot and Skempton coefficients govern HM rock behavior

- Effective stress $\Delta\sigma'$ controls deformation of saturated porous materials because part of the total load $\Delta\sigma$ acts on the fluid pressure Δp
- In **undrained conditions** (or in the short-time), an increase in load (total stress, $\Delta\sigma$) causes a proportional increase in fluid pressure Δp
- The product αB defines the impact of an applied load on the solid skeleton, and thus the material deformation, under undrained conditions.

$$\sigma' = \sigma - \alpha p$$

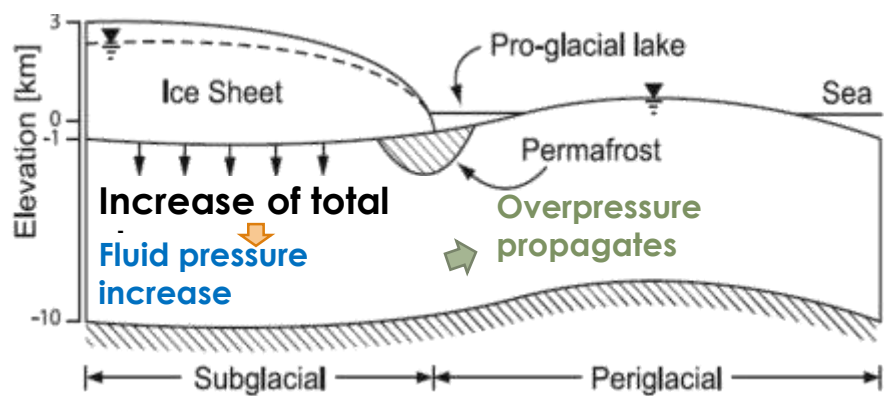
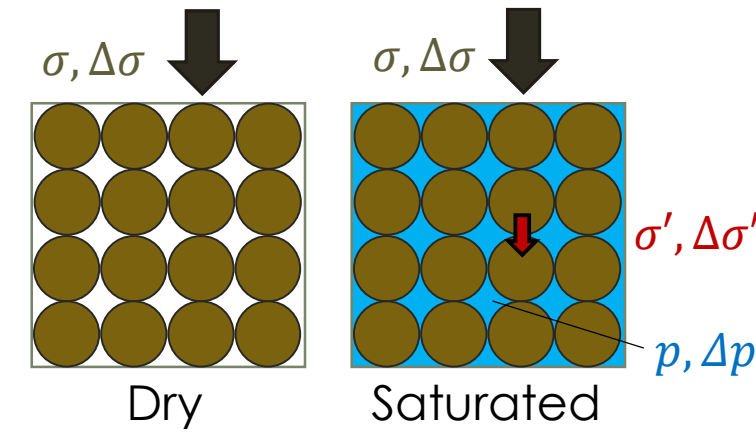
$$\Delta\sigma' = \Delta\sigma - \alpha \Delta p$$

Biot coefficient α

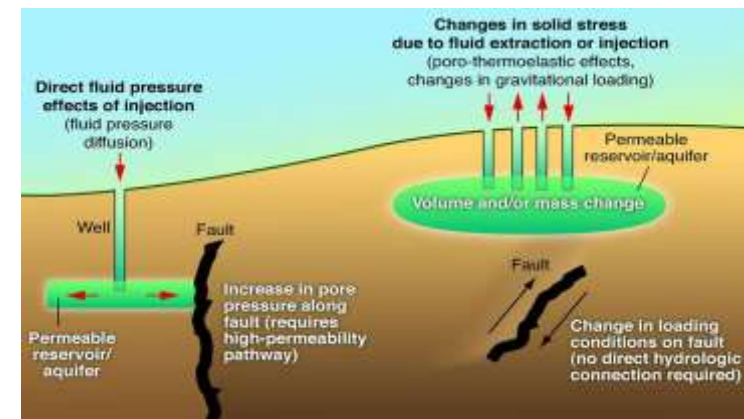
$$\Delta p = B \Delta\sigma$$

Skempton coefficient B

$$\Delta\sigma' = (1 - \alpha B) \Delta\sigma$$



Modified from Lemieux et al. (JGR, 2008)



Ellsworth (Science, 2013)

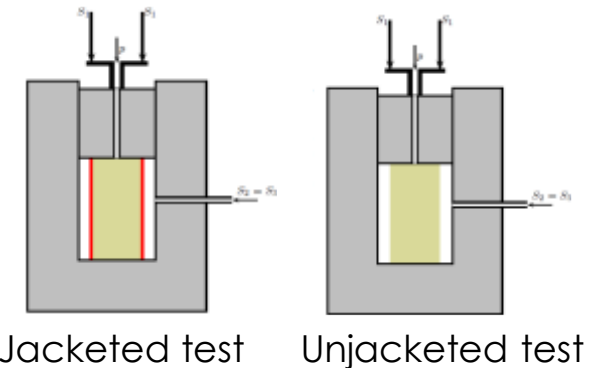
Some details about the poroelastic coefficients (*)

- Biot coefficient α reflects the effects of pore fluid pressure on the solid matrix
- It depends on the petrophysical properties of the solid skeleton at the microscale
- Skempton coefficient B defines the pressure variation in response to total stress variation under undrained conditions
- It depends on rock and fluid properties at bulk scale
- They range between 0 and 1, being ≈ 1 in highly compressible materials (e.g., soils), and < 1 in stiff formations (e.g., crystalline rocks);
- They can be directly or indirectly measured in the lab, or they can be calculated from known properties through theoretical expressions valid for homogeneous isotropic media

$$\sigma' = \sigma - \alpha p$$

$$\alpha = \left(\frac{\partial \sigma_m}{\partial p} \right)_V$$

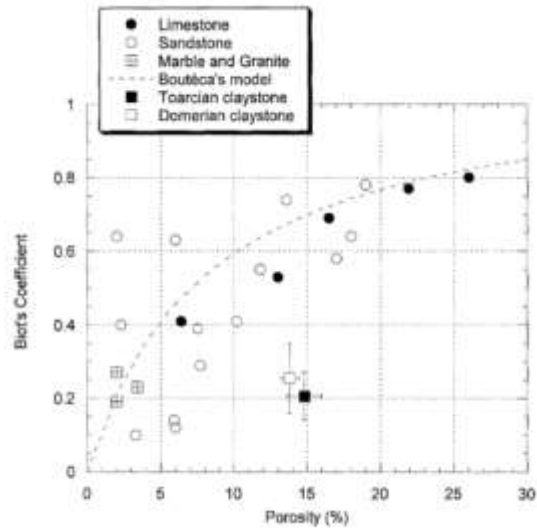
$$B = \left(\frac{\partial p}{\partial \sigma_m} \right)_\zeta$$



* Refs: Biot 1941; see also: Detournay & Cheng 1993; Cheng 2016; Coussy 2004

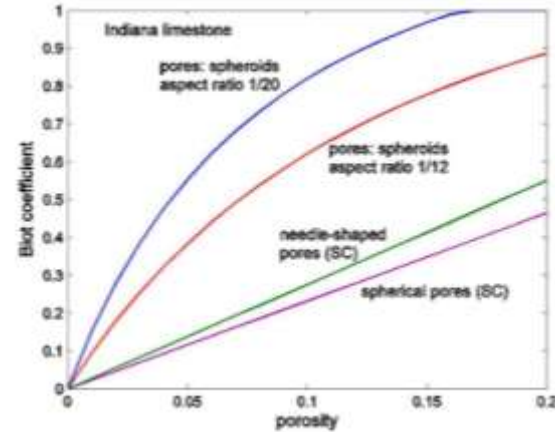
The coefficients largely depend on pores and their shape

α increases with **porosity**

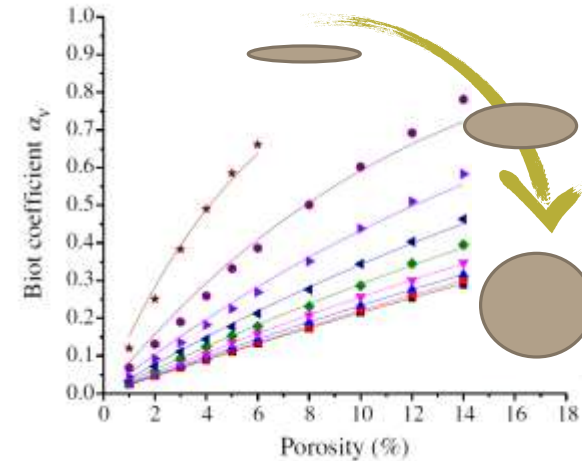


Experimental data
(Selvadurai, *Geosciences*, 2021)

α increases with **elongated** pores and cracks

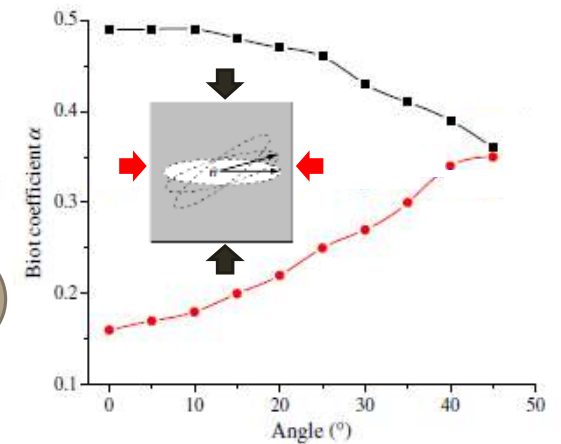


Theoretical estimations
(Selvadurai & Suvorov, *Sci. Rep.*, 2020)



Numerical estimations
(Modified from Tan & Konietzky, *Tecton.*, 2014)

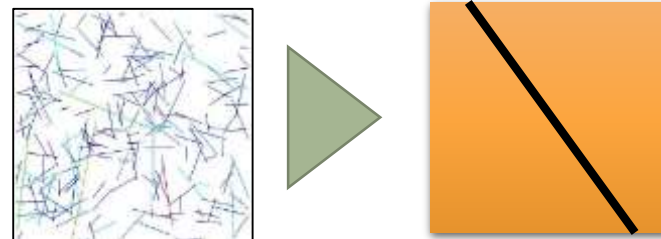
α is **anisotropic**



† studies limited to sample scale



What about large-scale fractured rocks?

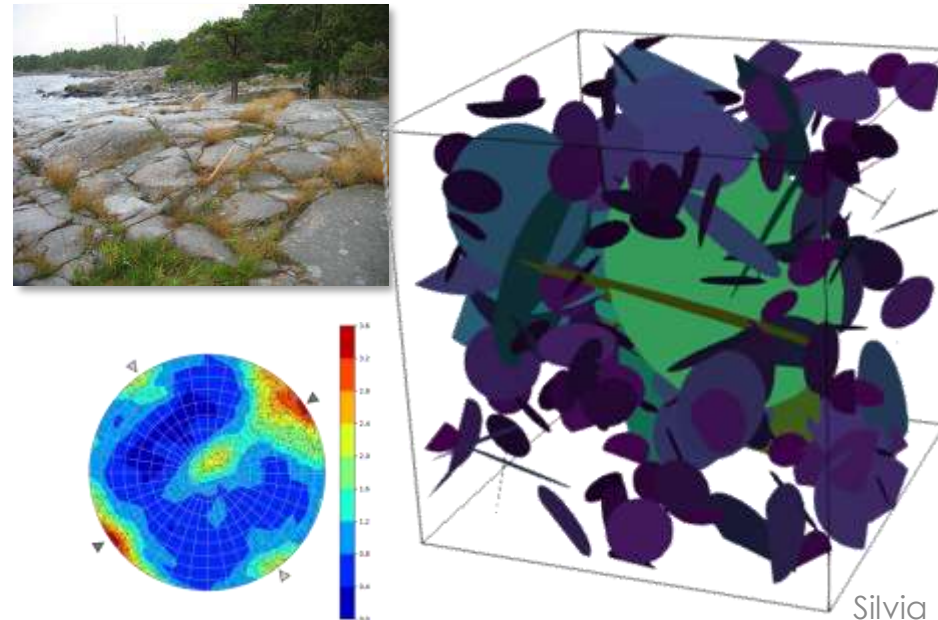


Field-scale problems require estimating effective poroelastic coefficients referring to the rock mass

Proposed estimations for fractured rocks:

- Poroelasticity theory for anisotropic porous media (e.g., Cheng 1997+ extensive literature)
- Equivalent stiffness for fractured rocks + traditional theoretical expressions (e.g., Wong 2017; Selvadurai et al. 2019; Selvadurai and Suvorov 2020; Berryman 2012)
- Numerical estimations from observed response (e.g., Chen et al. 2020)
- Volume-weighted averaging (e.g., Tuncay&Corapcioglu 1995; Tan&Konietzky 2014)

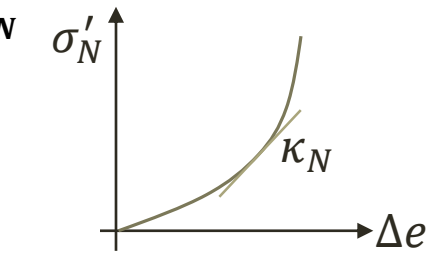
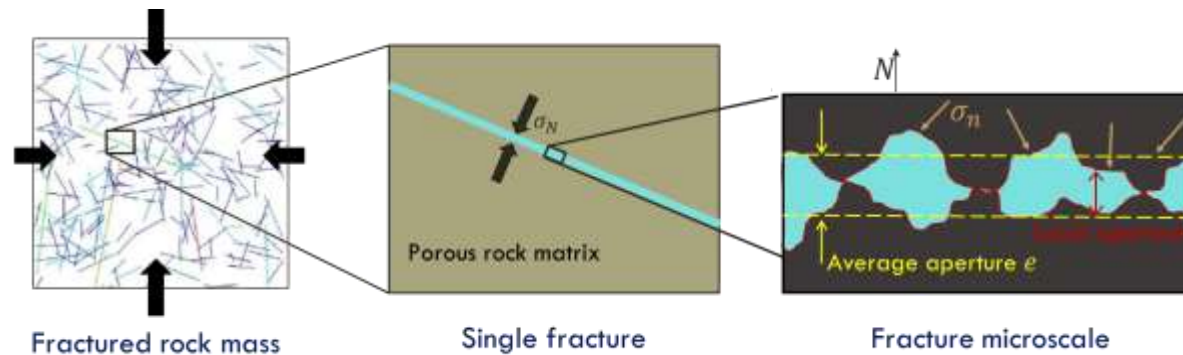
- ☹ Limited to sample scale
- ☹ Limited to 2D
- ☹ Not validated
- ☹ No consideration of fracture orientation
- ☹ No consideration of fracture volume
- ☹ No porosity of intact rock



Defining equivalent Biot and Skempton coefficients for a single saturated fracture

Assumptions:

- Fracture with aperture e is **fluid-filled**
- Fracture characteristics (aperture, stiffness) are **homogeneous in the plane**
- The resulting force is perpendicular to the average fracture plane (deviations due to asperities cancel out)
- Fracture behavior is locally linear elastic with **normal stiffness** κ_N



$\sigma_N = \text{normal stress}$

$$\Delta\sigma'_N = \Delta\sigma_N - \alpha^f \Delta p$$

$$\Delta p = B^f \Delta\sigma_N$$

Fracture Biot coefficient

- $\alpha^f = 1$ for open fractures, while $\alpha^f = 0$ for sealed fractures.

Fracture Skempton coefficient

- analytically derived considering a volumetric approach $B^f = (\beta e \kappa_N + \alpha_N)^{-1}$,
 $\beta = \text{fluid compressibility}$

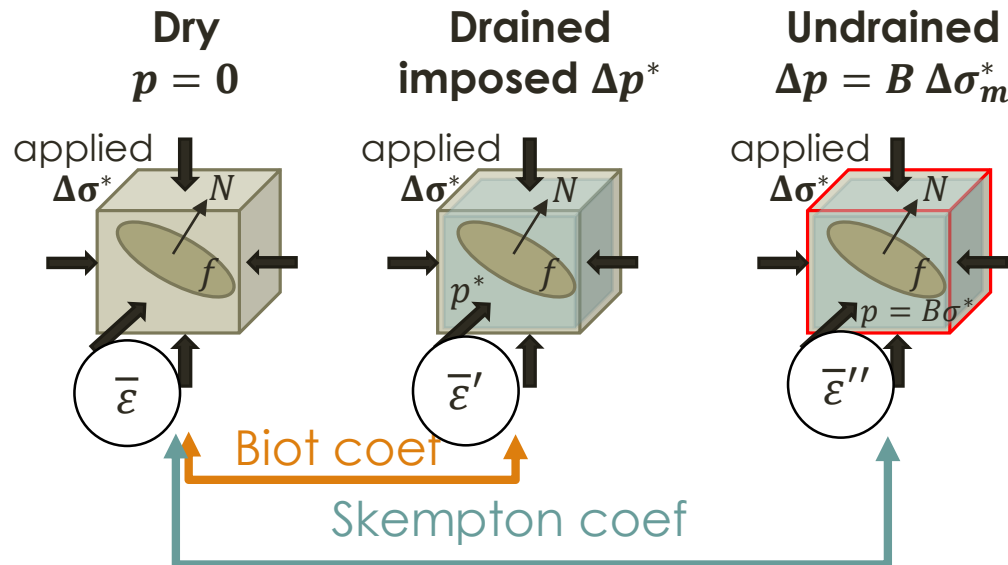
Conceptual model to define equivalent Biot and Skempton coefficients for a saturated fractured rock mass

$$\Delta\sigma'_m = \Delta\sigma_m - \bar{\alpha}\Delta p, \text{ equivalent Biot coeff.}$$

$$\Delta p = \bar{B}\Delta\sigma_m, \text{ equivalent Skempton coeff.}$$

$\Delta\sigma_m$ = average stress

- Equivalent coefficients are those that control the effective deformation.
- They are derived by comparing the **total volume deformations** under an applied incremental stress tensor $\Delta\sigma^*$ and different hydraulic conditions.



$$\frac{\bar{\varepsilon}}{\Delta\sigma_m^*} = \frac{\bar{\varepsilon}'}{\Delta\sigma_m^* - \bar{\alpha} \Delta p^*} \Rightarrow \bar{\alpha} = \frac{\bar{\varepsilon} - \bar{\varepsilon}'}{\bar{\varepsilon}} \frac{\Delta\sigma_m^*}{\Delta p^*}$$

$$\frac{\bar{\varepsilon}}{\Delta\sigma_m^*} = \frac{\bar{\varepsilon}''}{(1 - \bar{\alpha} \bar{B}) \Delta\sigma_m^*} \Rightarrow \bar{B} = \frac{\bar{\varepsilon} - \bar{\varepsilon}''}{\bar{\varepsilon}} \frac{1}{\bar{\alpha}}$$

Defining equivalent Biot and Skempton coefficients for a saturated fractured rock mass

$$\Delta\sigma'_m = \Delta\sigma_m - \bar{\alpha}\Delta p, \text{ Biot coeff.}$$

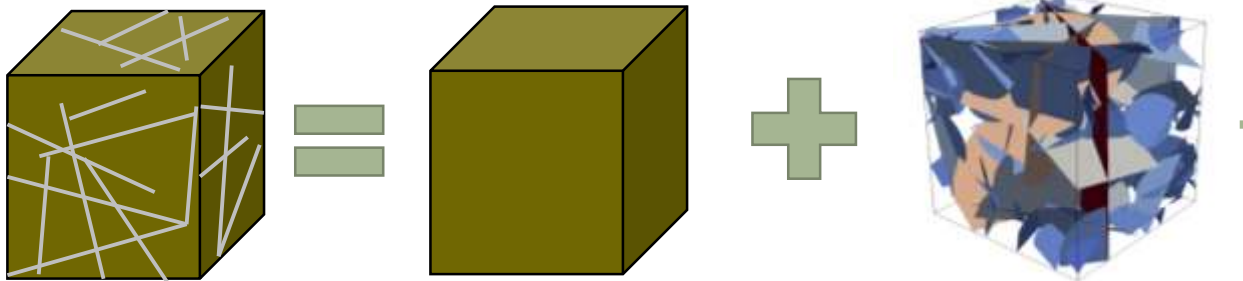
$$\Delta p = \bar{B}\Delta\sigma_m, \text{ Skempton coeff.}$$

$$\Delta\sigma_m = \text{average stress}$$

$$\bar{\alpha} = \frac{\bar{\varepsilon} - \bar{\varepsilon}'}{\bar{\varepsilon}} \frac{\Delta\sigma_m^*}{\Delta p^*}$$

$$\bar{B} = \frac{\bar{\varepsilon} - \bar{\varepsilon}''}{\bar{\varepsilon}} \frac{1}{\bar{\alpha}}$$

Total volume variation $\Delta V = \Delta V^r + \sum_f \Delta V^f$
 (non-interaction approximation - Grechka & Kachanov, 2006))

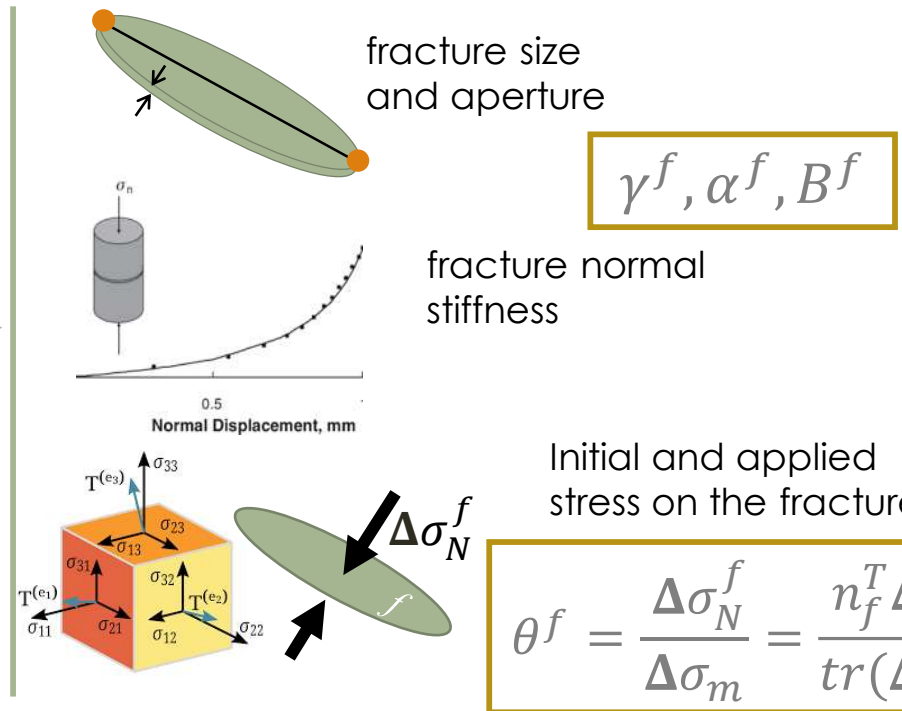


Intact rock (porous)

Fracture network

$$\gamma^r, \alpha^r, B^r$$

$\gamma^i = \Delta V^i / |\Delta\sigma^*|$
 contribution of element i to the total deformation

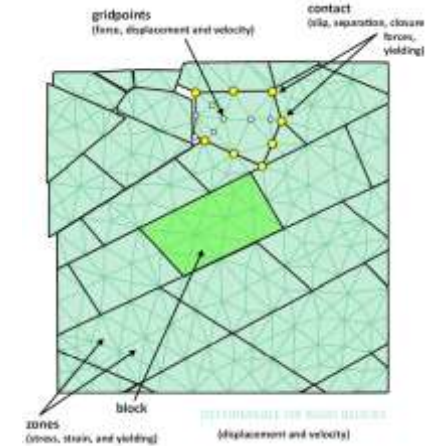
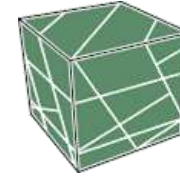


$$\bar{\alpha} = \frac{\gamma^r \alpha^r + \sum_f \gamma^f \alpha^f}{\gamma^r + \sum_f \theta^f \gamma^f} \quad \bullet \text{ } \gamma\text{-weighted averages}$$

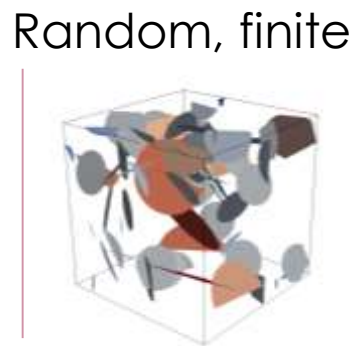
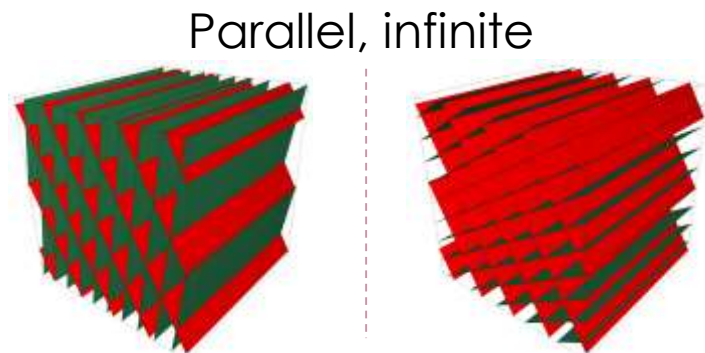
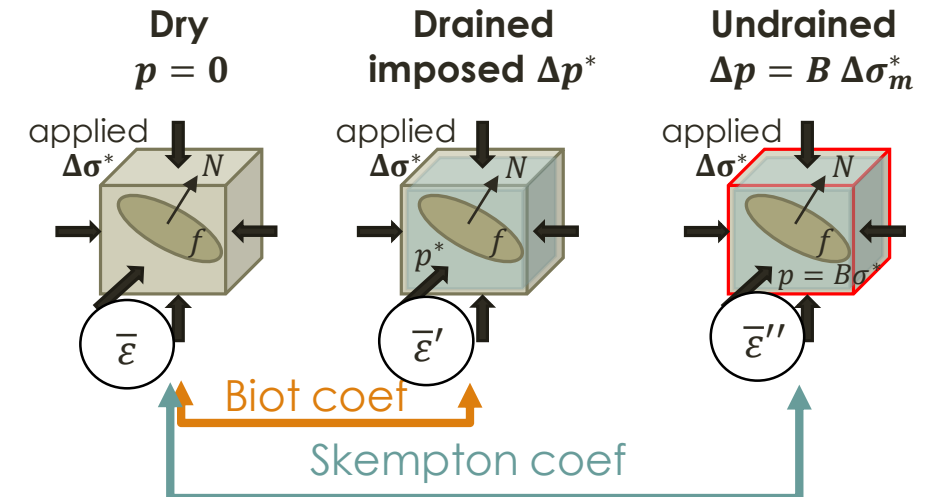
$$\bar{B} = \frac{\gamma^r \alpha^r B^r + \sum_f \gamma^f \theta^f \alpha^f B^f}{\gamma^r \alpha^r + \sum_f \gamma^f \alpha^f} \quad \bullet \text{ Depend on the load } \Delta\sigma$$

Validated against numerical results (3DEC)

- Fractures and rock blocks are explicitly represented, and their HM behavior simulated
 - Distinct Element Method
 - Assembly of deformable blocks limited by fracture planes
 - Elastic behavior inside fracture plane
 - Loaded by stress



- The three hydraulic conditions are reproduced, and the total volume variations estimated
- Different fracture settings
- Stress is alternatively applied in the 3 directions

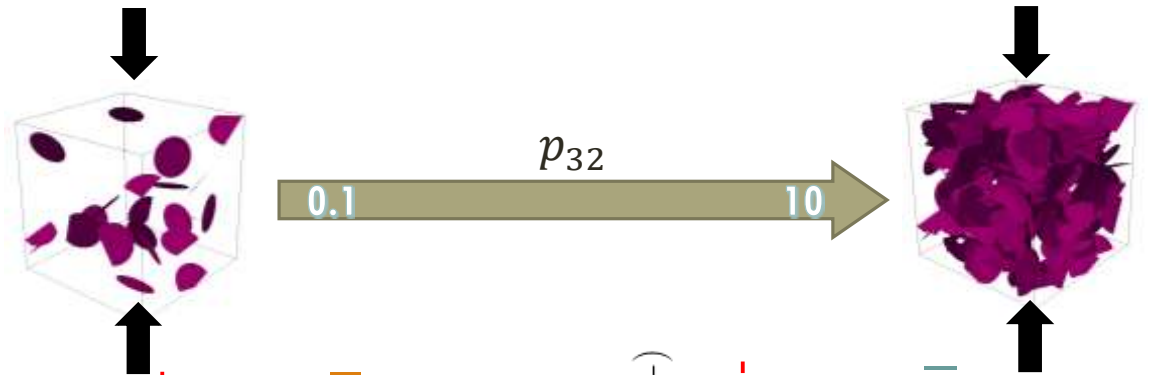


(De Simone et al., RMRE 2023)

Variability of equivalent coefficients when fracture aperture (e) and stiffness (κ_N) are constant (i)

$$\bar{\alpha} = \frac{\gamma^r \alpha^r + \sum_f \gamma^f \alpha^f}{\gamma^r + \sum_f \gamma^f} \quad \sum_f \gamma^f = \sum_f \frac{S^f}{\kappa_N} = p_{32} \kappa_{n,f}^{-1}$$

$$\bar{B} = \frac{\gamma^r \alpha^r B^r + \sum_f \gamma^f \theta^f \alpha^f B^f}{\gamma^r \alpha^r + \sum_f \gamma^f \alpha^f}$$

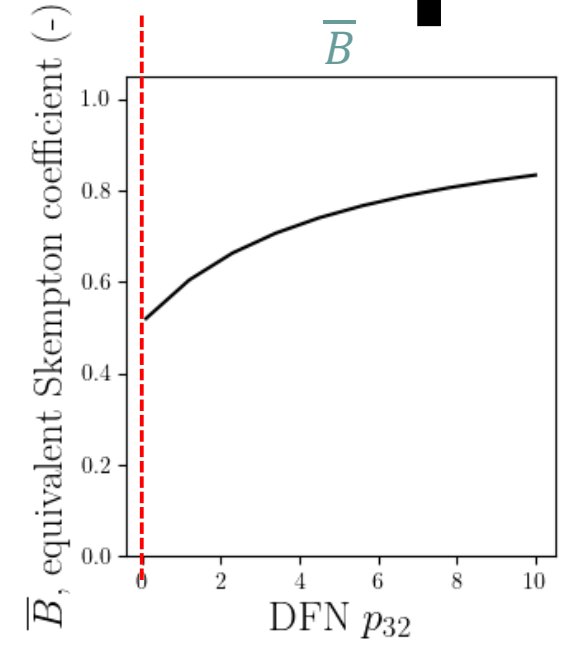
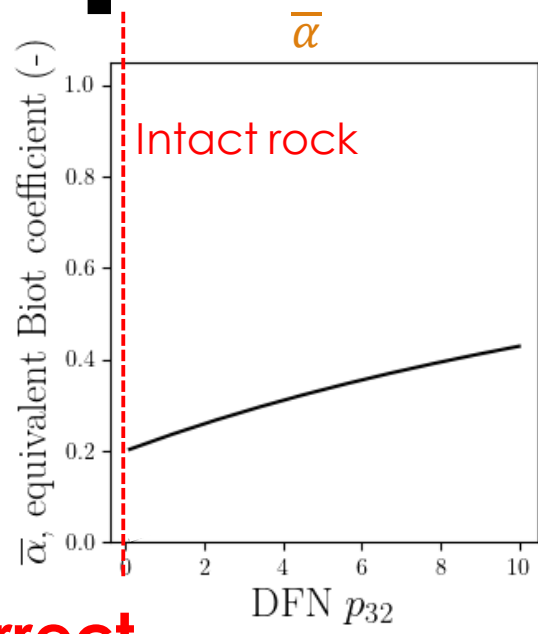


$\bar{\alpha}$ and \bar{B} increase with fracture density because fractures increase the overall rock mass deformability



smaller effective stresses $\Delta\sigma' \propto (1 - \bar{\alpha} \bar{B})$

smaller deformations?



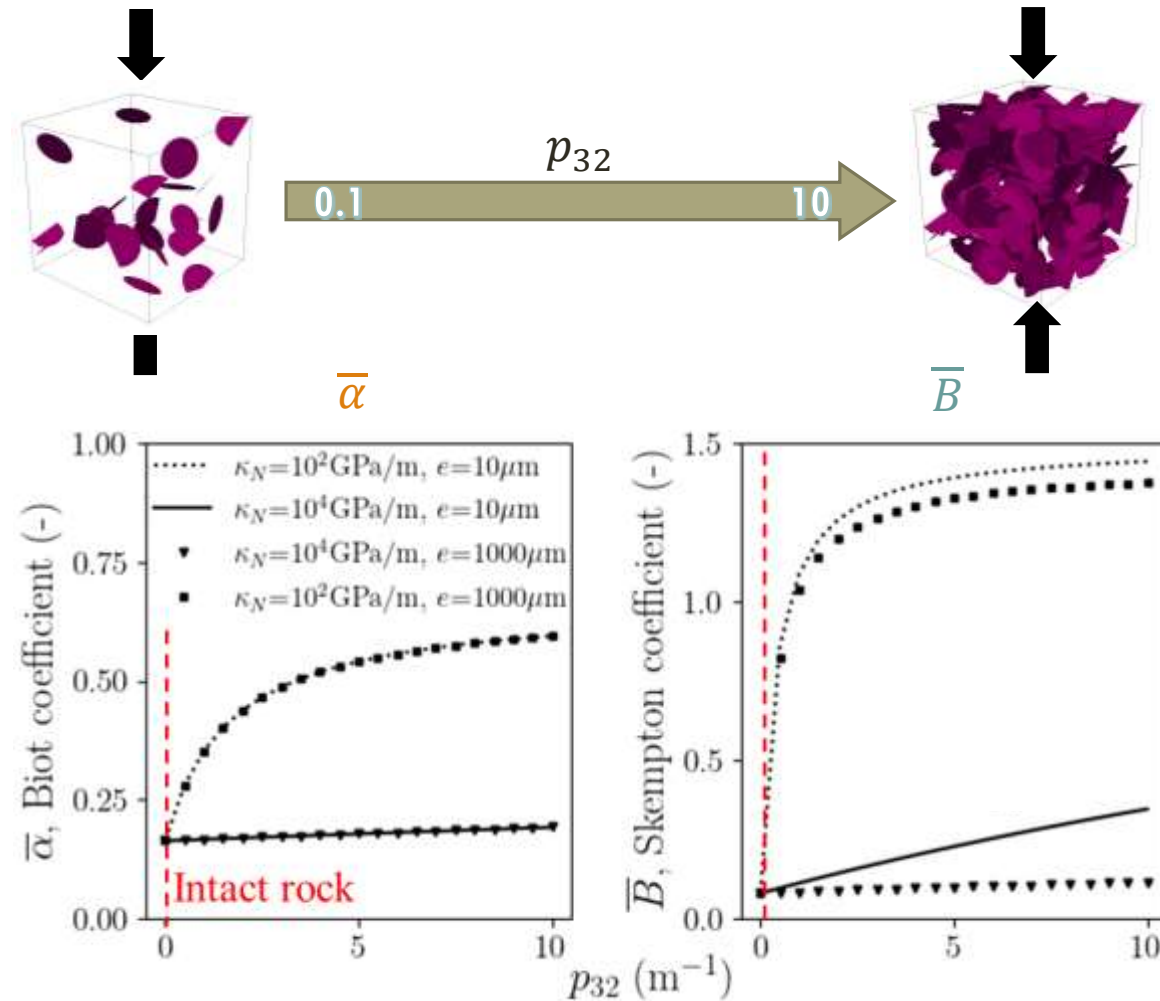
Neglecting fractures leads to incorrect estimations of the HM behavior

Variability of equivalent coefficients when fracture aperture (e) and stiffness (κ_N) are constant (ii)

$\bar{\alpha}$ and \bar{B} depend on fracture stiffness and aperture

- They both decrease with increasing fracture stiffness
- Limited effects of aperture – \bar{B} decreases with increasing fracture aperture

Uncertainty in characterizing the HM behavior



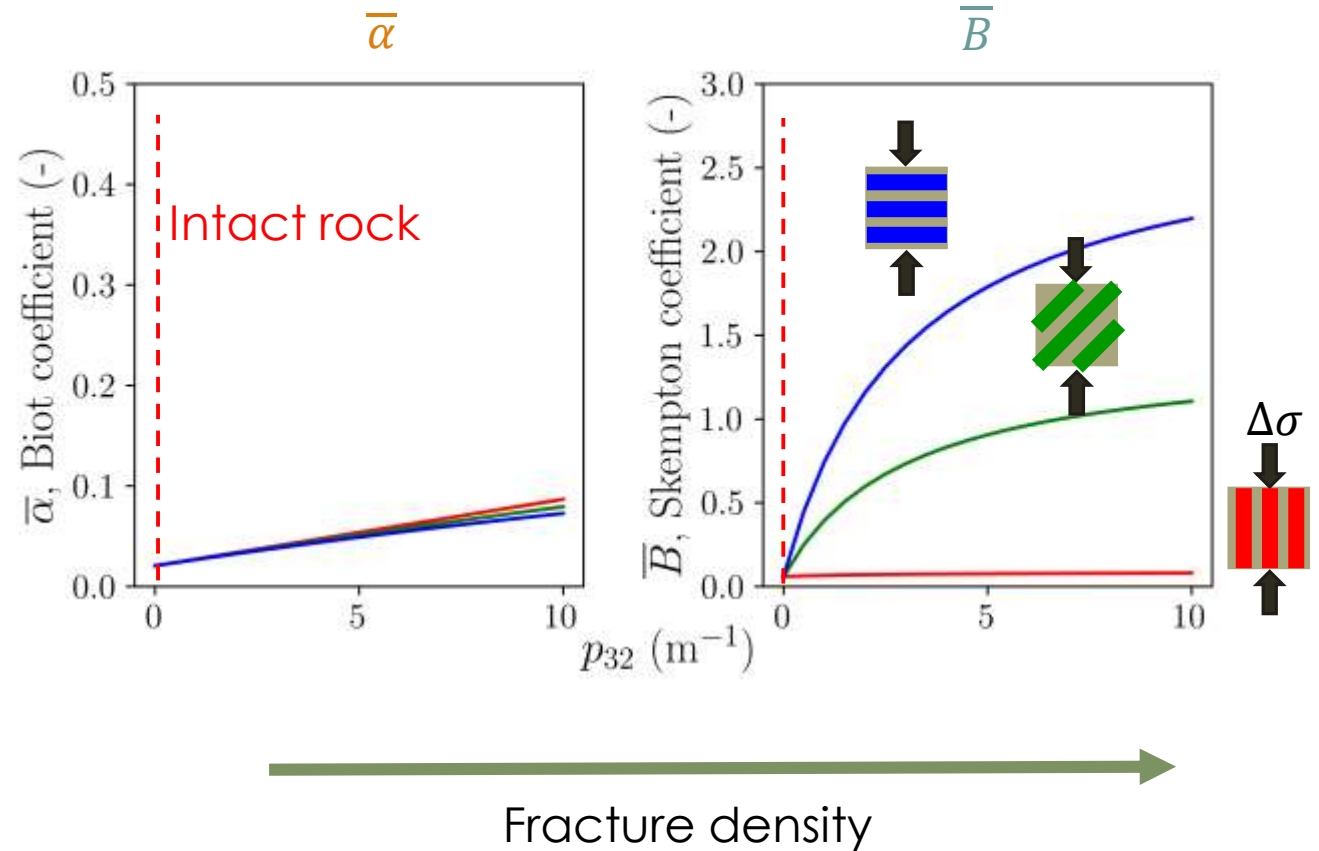
(De Simone *et al.*, RMRE 2023)

Variability of equivalent coefficients when fracture aperture (e) and stiffness (κ_N) are constant (iii)

$\bar{\alpha}$ and \bar{B} depend on fracture orientation with respect to applied stress because orientation affects fracture contribution to the overall deforming capacity

\bar{B} is larger when the applied stress acts normal to the fractures

The opposite for $\bar{\alpha}$ but very limited

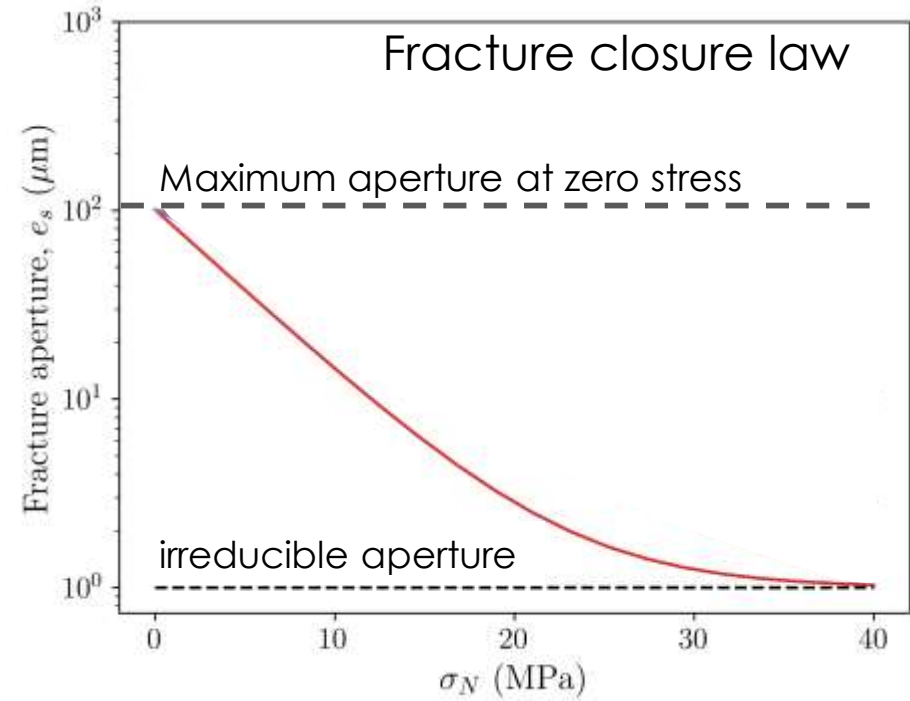


HM behavior is strongly anisotropic

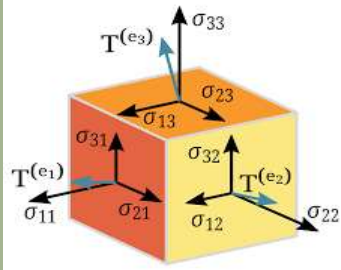
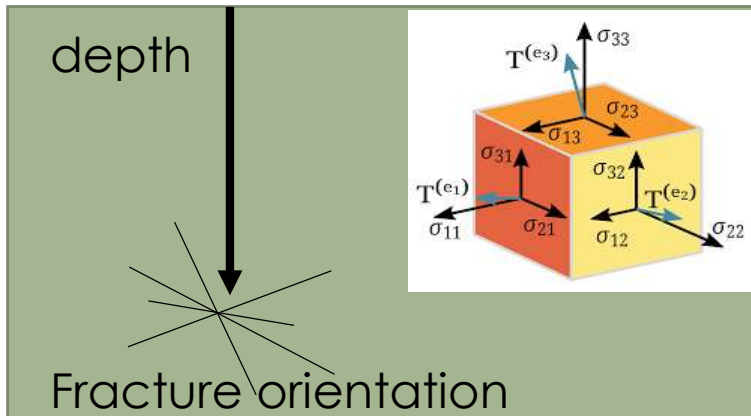
But fracture aperture (e) and stiffness (κ_N) are stress-dependent

Fracture aperture decreases **non-linearly** with increasing normal stress, because fracture stiffness increases as the normal stress increases

- hyperbolic (Goodman, 1976; Bandis et al., 1983)
- logarithmic (Evans et al., 1992)
- exponential law (Liu et al., 2012)



Fracture aperture and stiffness depend on:



In-situ stress tensor

$\Delta\sigma_N$
↑

aperture
↓

stiffness
↑

Particularly affects the coefficients (they decrease with increasing stiffness)

Fracture orientation plays two effects

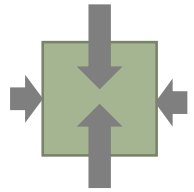
CONSTANT APERTURE AND STIFFNESS

- $\bar{\alpha}$ and \bar{B} depend on fracture **orientation** with respect to **applied stress** $\Delta\sigma$

STRESS-DEPENDENT APERTURE AND STIFFNESS

- $\bar{\alpha}$ and \bar{B} depend on fracture **orientation** with respect to **initial stress**

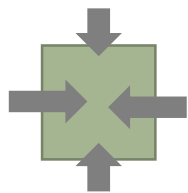
Initial stress



$\bar{\alpha}$: The effect of the orientation on $\Delta\sigma_N$ is **amplified** by its effect on κ_N

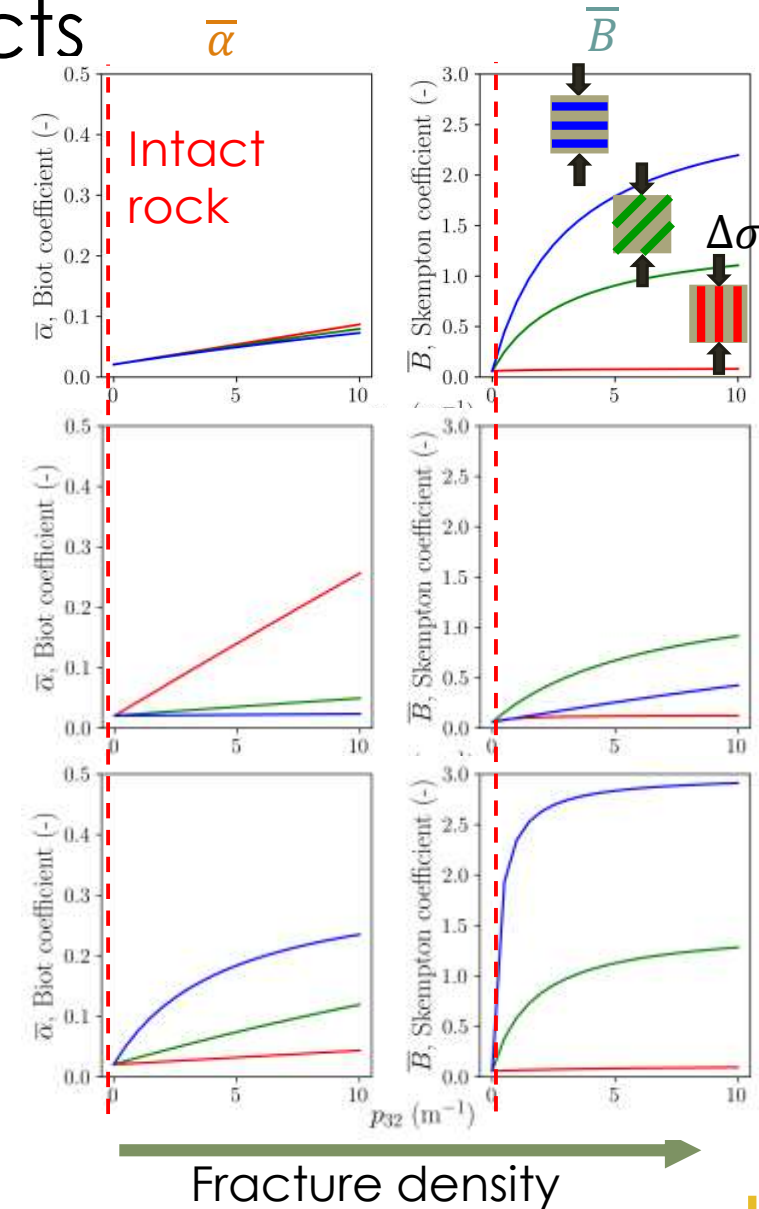
\bar{B} : The effect of the orientation on $\Delta\sigma_N$ is **counterbalanced** by its effect on κ_N

Initial stress

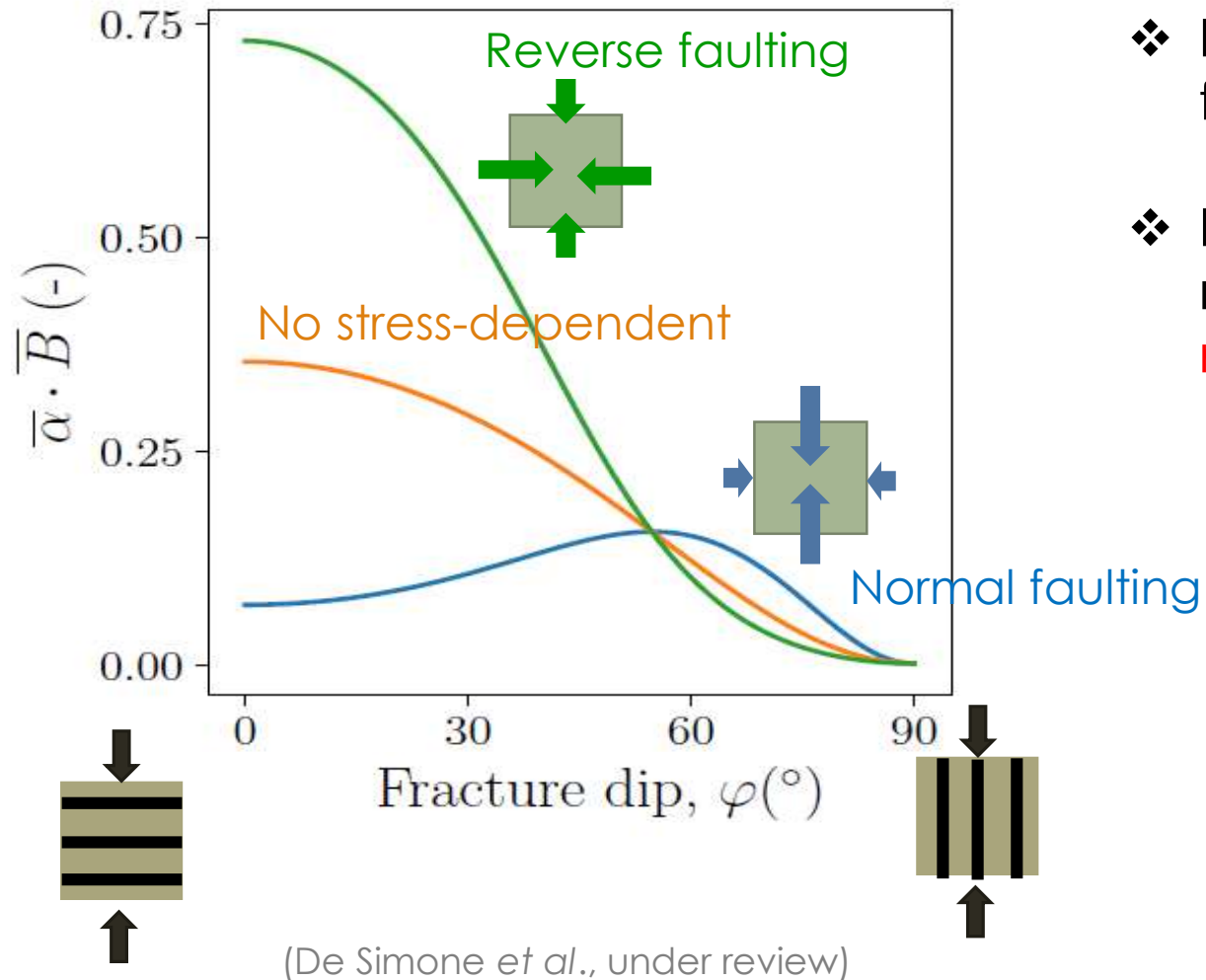


The opposite

HM behavior is even more anisotropic



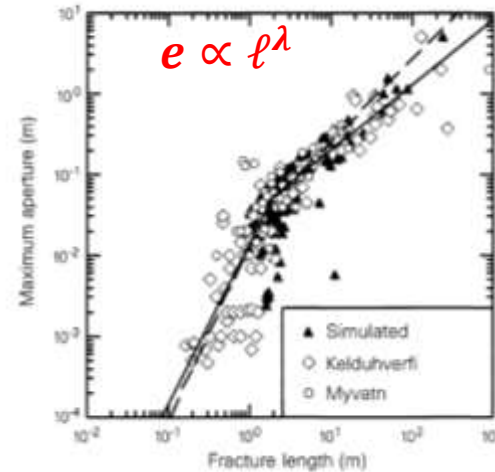
In undrained conditions, the effective stress is minimized for:



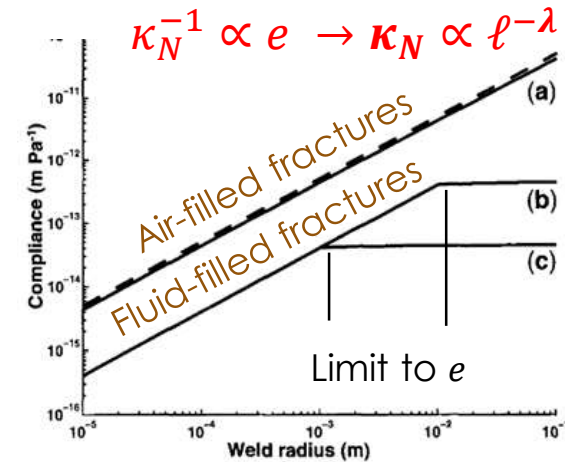
- ❖ Horizontal fractures under reverse faulting
- ❖ Fractures with dip=60° under normal faulting (which are the most critical for failure!)

Fracture aperture (e) and stiffness (κ_N) are size-dependent

Large fractures are more compliant and open

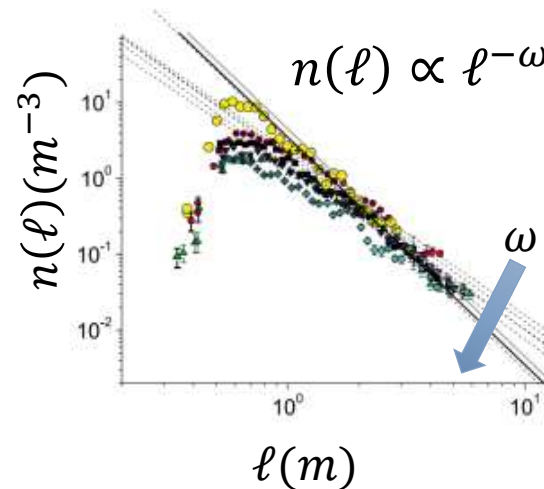
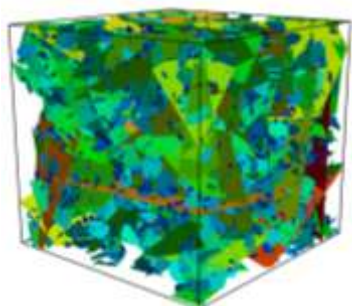


(Renshaw & Park, *Nature*, 1997)



(Worthington, *GSL*, 2007)

The distribution of fracture size follows a power law



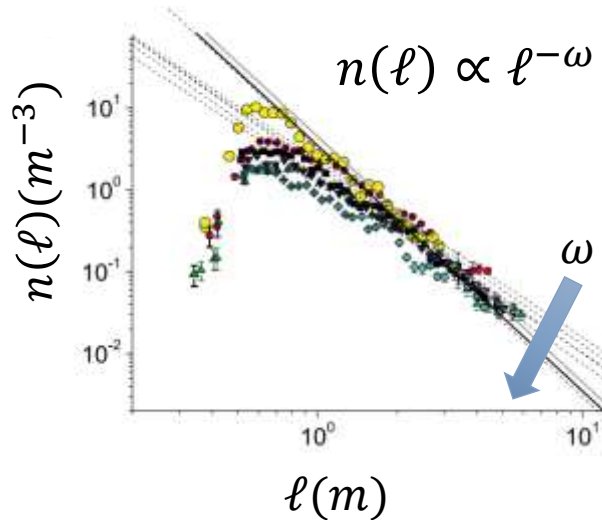
$$\sum_f \gamma^f = \sum_f \frac{S_f}{\kappa_N^f} \propto \int l^{2+\lambda} n(l) dl$$

Size distribution impacts the coefficients

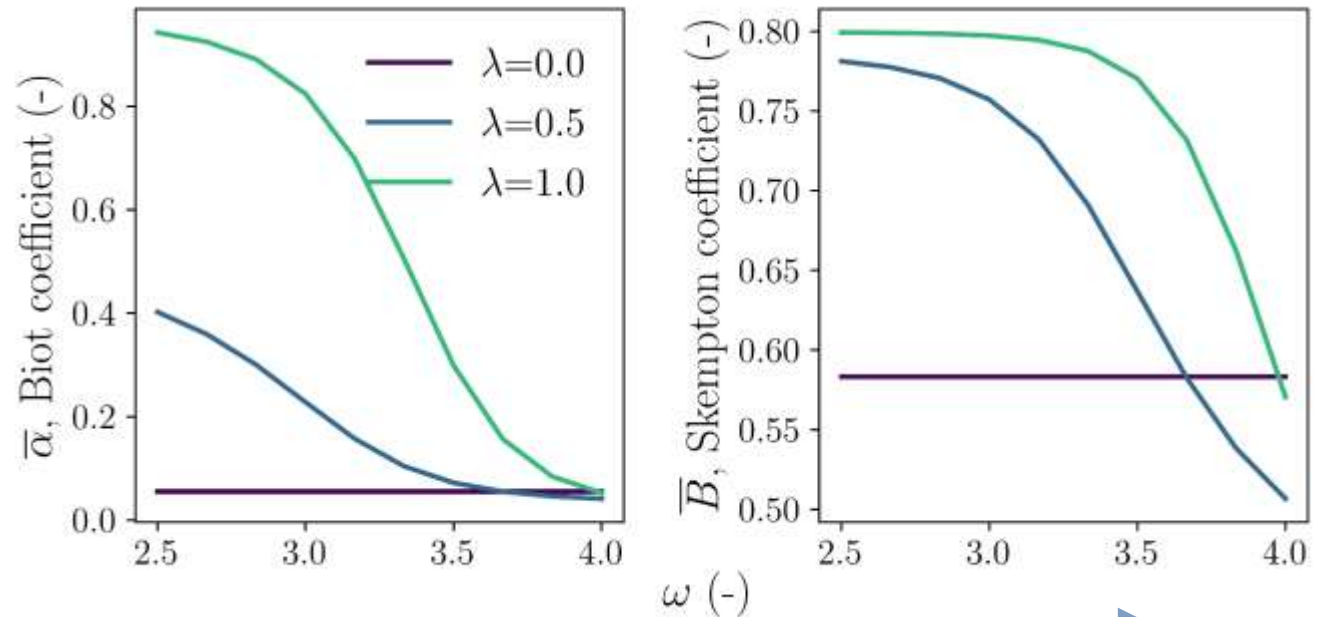
Assuming all fractures have the same orientation

$$\kappa_N \propto \ell^{-\lambda}$$

$$e \propto \ell^\lambda$$



$\bar{\alpha}$ and \bar{B} larger in systems populated by a few large fractures than in systems with many small fractures (for equivalent p_{32})



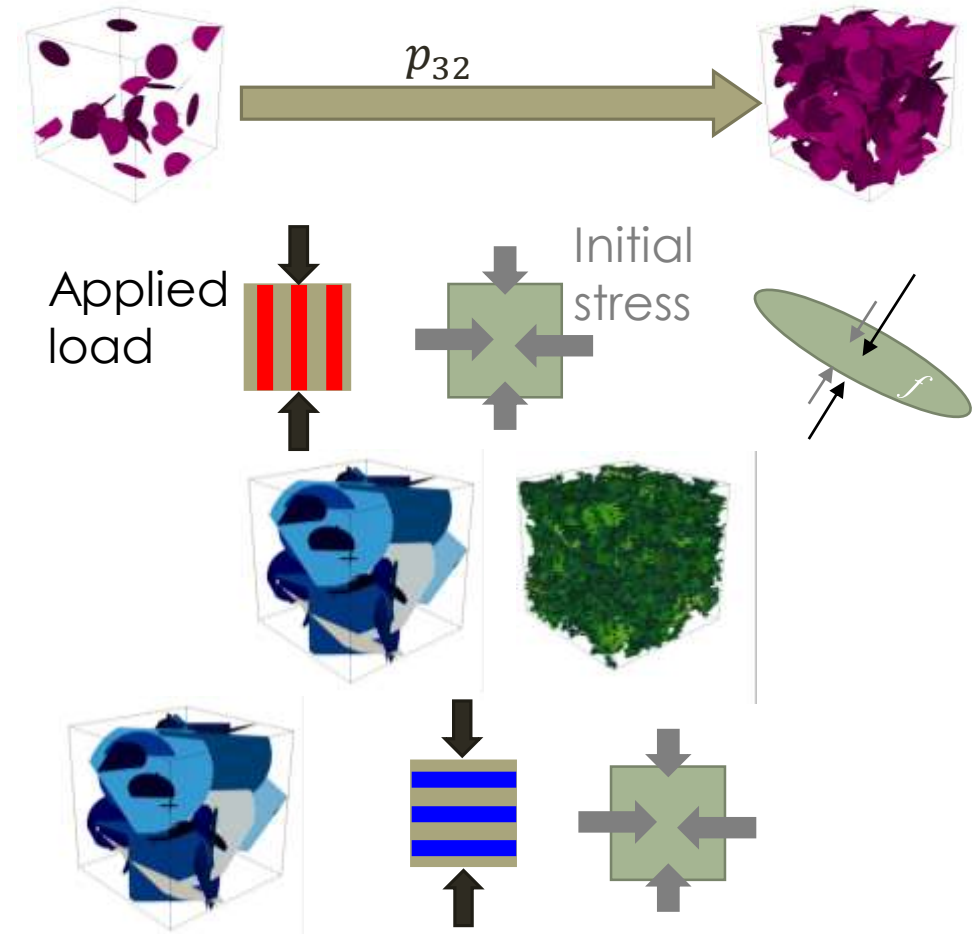
More small fractures

HM behavior determined by fracture size organization

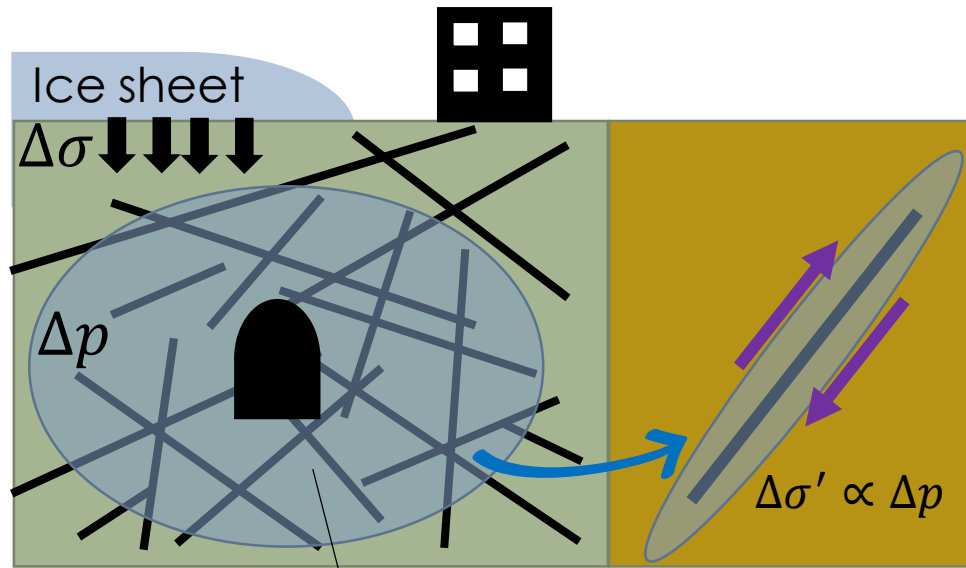
(De Simone *et al.*, under review)

Take-home messages

- Disregarding fractures lead to incorrect predictions: both coefficients increase with fracture **density**
- Strong **anisotropy**: Biot and Skempton coefficients depend on the **orientation** of the applied load, the initial stress and the fracture orientations
- Impact of the distribution of **fracture size**: Biot and Skempton coefficients are larger in systems populated by a few large fractures
- The fracture contribution is larger in systems containing **large fractures oriented parallel to the largest principal initial stress and normal to the applied stress**
- **Uncertainties** arise from parameters that are not yet constrained by measurements



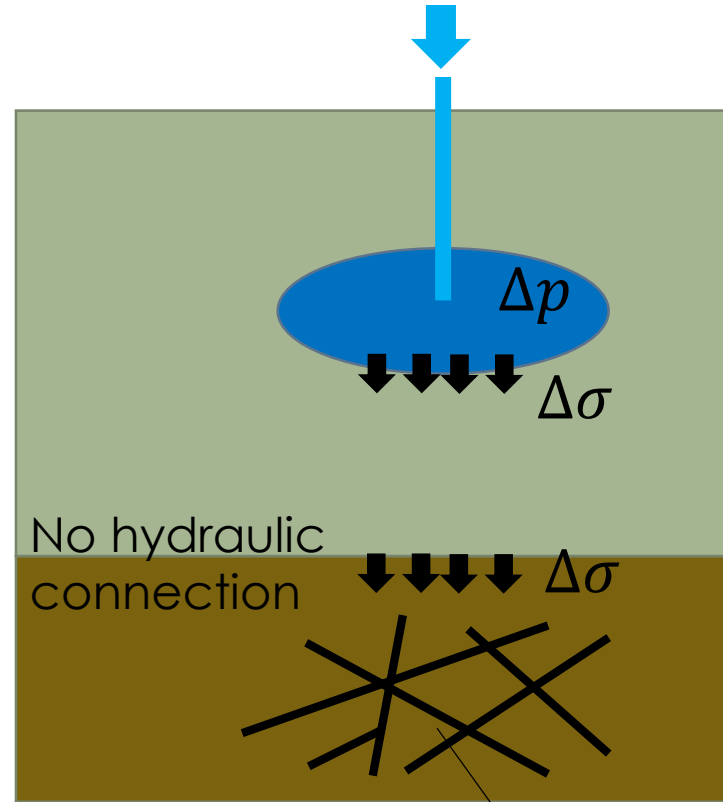
Implications for HM behavior and induced seismicity



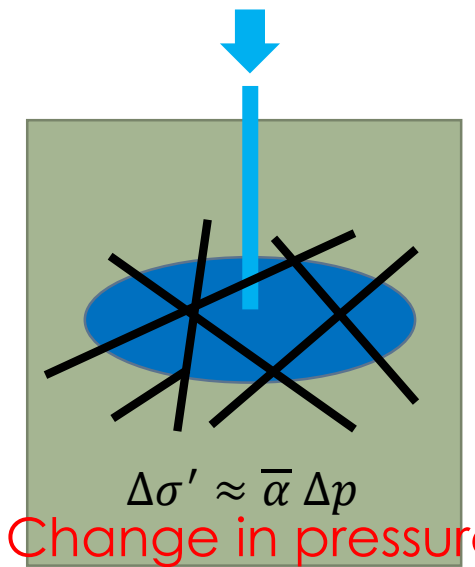
Change in pressure

$$\Delta\sigma' = (1 - \bar{\alpha} \bar{B}) \Delta\sigma$$

Change in loading conditions

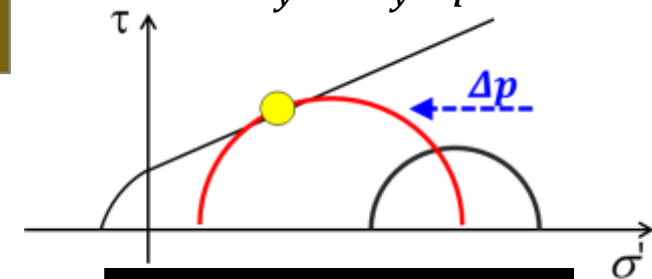


Change in loading conditions



$$\Delta\sigma'_x \approx \bar{\alpha}_x \Delta p$$

$$\Delta\sigma'_y \approx \bar{\alpha}_y \Delta p$$



Stay tuned!



Thank you for your attention!

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